

CLIMATIC OSCILLATIONS MODULATING THE LATE HOLOCENE FLUVIAL DISCHARGE AND TERRIGENOUS MATERIAL SUPPLY FROM THE RÍO DE LA PLATA INTO THE SOUTHWESTERN ATLANTIC OCEAN

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Abstract

The Río de la Plata estuary (RdlP) and adjacent continental shelf exhibit complex hydrological processes as a result of a highly variable fluvial discharge and associated terrigenous supply, which are primarily controlled by regional climatic forcing in interaction with the oceanographic system. Previous hydrological studies indicate that Little Ice Age and Medieval Climate Anomaly resulted in increased/decreased river runoff due to an enhanced/weakened precipitation regime. This study aims to analyze the cyclicity in the runoffrelated continental supply and the associated climatic oscillation behind it over the past 1,000 cal yr BP. To achieve this, we used a 10-m-long sediment core retrieved from the RdlP mud depocenter, which exhibits an extremely high and fairly constant linear sedimentation rate of 1.1 cm yr⁻¹. We performed continuous 1-cm XRF element intensity scans and performed time series analysis on Fe/K, Ti/Al, Fe/Ca, Ti/Ca and Si/Al ratios, considered as regional proxies for inferring fluvial supply. The most significant cyclicities occur at 215, 192, 115, 100, 49, 47, 37, 35, and 2.5-8 yrs recurrence times. The long-term cyclicity is persistent throughout the

1. Introduction

Climate variability exerts a direct influence on regional precipitation and continental freshwater supply (Barreiro,

whole record for all ratios, and it is probably related to solar forcing, i.e., the 200 yr Vries/Suess cycle, which determines the activity of the South American Monsoon System activity. The shorter climate-related cyclicities (<100 yr) are probably related to the Atlantic Multidecadal Oscillation, the Pacific Decadal Oscillation and the El Niño Southern Oscillation climatic modes, and were registered only for Ti/Al, Fe/K and Si/Al ratios. We further identified an intensification of such shorter climate-related cyclicities over the past 500 yr BP. This study provides new evidence that both tropical Pacific and Atlantic SSTs and solar forcing are the main drivers of changes in rainfall over Southeastern South America on interannual to decadal, multidecadal and centennial scales, and thus modulates the RdIP river discharge.

Keywords: Climatic variability. Spectral analyses. XRF element ratios. Mud depocenter. Continental shelf.

2010; García-Rodríguez et al., 2014, Lüning et al., 2018), on oceanographic processes, and also on sediment transport, resuspension and deposition from the continent to the continental shelf (Lantzsch et al., 2014; Marrero et al., 2014;



RESEARCH PAPER

Briceño-Zuluaga et al., 2016; Perez et al., 2016). Large estuaries are responsible for most of the terrigenous material supply into the continental margins, and the associated sediment transport processes mobilize terrigenous chemical elements to the coastal zone and continental margins (Depetris, 1968; Bianchi et al., 2007; García-Rodríguez et al., 2014).

In this sense, the chemical elements Ti, Fe and Al are mainly related to aluminosilicates and namely clay minerals as weathering products from the continent (Goldberg and Arrhenius, 1958; Jansen et al., 1998; Yarincik et al., 2000). Therefore, such elements reflect the terrigenous contribution in coastal and offshore sediments (Martins et al., 2007; Burone et al., 2013). The chemical element K in marine sediments is associated with terrigenous material (Goldberg and Arrhenius, 1958). This chemical element is relatively abundant in arid regions where chemical weathering rates are usually low (Govin et al., 2012). Si is an important nutrient in the oceans as it is a basic component of the skeletons of diatoms, radiolarians, silicoflagellates and sponges (DeMaster, 2002; Libes, 2009). Thus, temporal and spatial variability in the annual cycles of Si availability in estuaries and coastal areas depend on both river supply and biological production (Bianchi, 2007). In contrast, Ca mainly reflects the marine biogenic carbonate content in the sediment, and is thus associated with local marine productivity (Haug et al., 2001; Salazar et al., 2004; González-Mora and Sierro, 2007).

Govin et al. (2012) assessed the potential use of Ti/Ca, Fe/Ca, Ti/Al, Fe/K and Si/Al element ratios for Southwestern Atlantic Ocean (SWAO) sediments, to reconstruct different terrestrial past climate regimes. These authors demonstrated that Ti/Ca and Fe/Ca resemble the terrigenous supply versus the marine influence, as most of the Ca corresponded to carbonates, related to a marine origin. Furthermore, the Ti/Al ratio is appropriate to infer the variability in the input of aeolian versus fluvial terrigenous material, in areas of dust deposition. Also, the Fe/K and Si/Al are suitable to reconstruct South American climatic zones, with high values of the Fe and Al, registered in highly weathered tropical areas, and high values of K and Si, in slightly weathered drier regions, and thus are both proxies of the continental runoff (Govin et al., 2012).

Seager et al. (2010) determined that the variability in sea surface temperature (SST) in both the tropical Pacific and Atlantic Oceans cause most of the Southeastern South American (SESA) temporal precipitation changes. These fluctuations account for most of the interannual to multidecadal variability, with SST in the tropical Pacific dominating at interannual timescales, and those of the tropical Atlantic, dominating at long-term scales (Seager et al., 2010). The interannual variability (i.e., 2-8 yrs. cycles) is related to El Niño Southern Oscillation (ENSO; Depetris and Pasquini, 2007b; Garreaud et al., 2009). The decadal and multidecadal climatic variability in SESA mainly results from two competing climatic modes: the Pacific Decadal Oscillation (PDO) with 30 yr. cycles (Garreaud et al., 2009) and the Atlantic Multidecadal Oscillation (AMO) with 64 yr. cycles (Delworth and Mann, 2000).

In the Río de la Plata (RdlP) watershed, the seasonal variability in rainfall is mainly linked to a latitudinal migration of the Intertropical Confluence Zone (ITCZ), an equatorial band of convective activity that migrates northwards in winter, and southwards in summer forming the central element of the South American Monsoon System (SAMS; Zhou and Lau, 1998). During the mature phase of SAMS in summer, intensive convective activity is observed (Fig. 1), associated with an enhancement of the equatorial North Atlantic trade winds, a buildup of strong Northwesterlies along the eastern side of the tropical Andes, and a migration of the South Atlantic Convergence Zone (SACZ) towards its southernmost position (Zhou and Lau, 1998). The SACZ is related to a convective band that extends from the Amazonas Basin to the SWAO (Carvalho et al., 2004; Garcia and Kayano, 2010). Hence, SAMS activity is strongest during austral summer, and SACZ leads to an increase in rainfall over the RdlP watershed (Fig. 1; Robertson and Mechoso, 2000; Chiessi et al., 2009). The opposite pattern is commonly observed during austral winter when SAMS activity is weakest (Fig. 1; Carvalho et al., 2004; Chiessi et al., 2009).

There is significant evidence on the climatic changes occurred during the late Holocene: i.e., the Medieval Climate Anomaly (MCA; AD 800–1300) and the Little Ice Age (LIA; AD 1400-1800) in SESA. Both MCA and LIA exerted an influence on the regional rainfall pattern, with distinct regional differences (Cioccale, 1999; Iriondo, 1999; Meyer and Wagner, 2009; Moy et al., 2009; Piovano et al., 2009; Bird et al., 2011; del Puerto et al., 2011, 2013; Vuille et al., 2012; Apaéstegui et al., 2014; Salvatteci et al., 2014; Perez et al., 2016). Several studies reported a decrease in SAMS activity during MCA and strengthening of SAMS during LIA when compared to pre-MCA times (Fig. 2; Bird et al., 2011; Vuille et al., 2012; Apaéstegui et al., 2014; Salvatteci et al., 2014). Such studies demonstrated that SAMS is sensitive to the latitudinal displacement of the ITCZ, as a response to temperature changes in the northern hemisphere, particularly over the North Atlantic Ocean (Bird et al., 2011). Thus, SAMS variability appears to be closely related to AMO (Chiessi et al., 2009; Apaéstegui et al. 2014). The strengthening/weakening of SAMS during the LIA/MCA led to an increase/decrease in rainfall over the RdlP watershed, with a concomitant increase/decrease in continental freshwater supply to the adjacent SWAO inner continental shelf (Fig. 2; Perez et al., 2016).

To our knowledge, high-resolution studies dealing with climatic oscillations and the associated changes in RdlP fluvial supply to the inner SESA shelf have not yet been undertaken. This study aims to address the late Holocene interannual, decadal, multidecadal and centennial cyclicities of RdlP fluvial discharge and to relate such variability to the



particular climatic oscillation modes. We analyzed a continuous dataset of high-resolution XRF element ratios using time series analyses to determine the related ciclicity in continental freshwater supply. The additional combination with biogenic Si data allows discriminating between the

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> particular biogenic and terrigenous Si contribution to the sedimentary record. The results significantly help to understand the long-term paleoenvironmental processes in highly dynamic shelf systems, under the regime of strongly coupled climatic-hydrologic-oceanographic interactions.



Fig. 1. Seasonal phases (mature and demise) of SAMS and the resulting seasonal precipitation pattern over the RdlP drainage basin (black dotted box). Source: http://www.cpc.ncep.noaa.gov/products/Global_Monsoons/American_Monsoons/SAMS-Phases/

1.1 Study area

The RdlP forms through the confluence of the Paraná and Uruguay Rivers and drains into the SWAO as the second largest fluvial system in South America, only exceeded by the Amazon River, at North Brazil (Acha et al., 2003). The watershed covers 20% of the continent (Acha et al., 2004, 2008). The Paraná River is the largest affluent of the RdlP accounting for 73% of its water discharge (Depetris and Paquini, 2007a). The RdlP exhibits maximum water discharge during summer, which, together with a dominance of NE winds, displaces the RdlP plume in the offshore direction. In winter, the river discharge decreases, coinciding with a dominance of SW winds that push the RdlP river plume northward and against the coast (Depetris and Paquini, 2007a; Piola et al., 2008). Most of the RdlP freshwater are supplied by the tropical upper Paraná drainage basin, but the total suspended sediment load is mainly supplied by the arid, mountainous and mostly sedimentary upper Bermejo River drainage basin (Depetris et al., 2003). A total of 129 x 10⁶ metric tons of suspended sediments is transported from the Paraná River to the SWAO (Depetris and Giffin, 1968).

The oceanographic system, i.e. the RdlP estuary and adjacent continental shelf, displays a pronounced salinity gradient due to the freshwater supply from the RdlP system, and, to a less extent, from the slightly northward located Patos Lagoon (Acha et al., 2004; Möller et al., 2008; Piola et al., 2008). Below 50 m water depth, oceanographic fronts and mixing processes of water masses of different salinity, nutrient concentration, and temperature are observed as a consequence of the highly energetic confluence of the shelf branches of Brazil and Malvinas Currents. An oceanographic shelf front extends across the outer shelf, i.e., the subtropical shelf front (Piola et al., 2005). Thus, the interaction of the RdlP and the SWAO forms a complex and dynamic system (Calliari et al., 2009).

The Uruguayan inner shelf hosts a high accumulation mud depocenter (Martins and Urien, 2004; Lantzsch et al., 2014; Perez et al., 2016). The depocenter forms below the 50 m isobath, inside the RdIP paleo-valley (promoting the deposition and avoiding the resuspension from strong shelf currents), which incised during the lowered sea level associated with the wide, Last Glacial Maximum shelf exposure (Masello and Menafra, 1998; Martins et al., 2003; Lantzsch et al., 2014).



RESEARCH PAPER

2. Materials and Methods

A 1028-cm long gravity core (GeoB13813-4) was retrieved from the RdlP mud depocenter (34°44.22'S, 53°33.27'W; 57 m water depth) during the research cruise M78/3a with the German research vessel *Meteor* in May 2009 (Krastel et al., 2012; Fig. 3a). During this expedition, subbottom echosounder profiling with the shipboard PARASOUND P70 system (4 kHz secondary parametric signal) showed an elongated depression on the seafloor corresponding to the RdlP paleo-valley, which is filled with a differentiated succession of acoustic facies (Fig. 3b; Krastel et al., 2012; Lantzsch et al., 2014).

2.1 Age-depth model

Six bivalve shells, evenly distributed over the core and preserved in life position, were collected for radiocarbon dating (Fig. 3c; Lantzsch et al., 2014; Perez et al., 2016). The samples were analyzed using AMS-¹⁴C (accelerated mass spectrometry) at the Poznan Radiocarbon Laboratory (Poland), (Tab. 1). The raw radiocarbon dates were calibrated using the calibration curve Marine13 (Reimer et al., 2013). The standard reservoir age of 324 ± 30 yr was applied during calibration data (southeastern Brazil; Alves et al., 2016). The age-depth model was developed with the Bacon software (v. 2.2; Blaauw and Christen, 2011; Tab. 1).



Fig. 2. Comparisons between late Holocene SESA records: SAMS activity changes (Bird et al., 2011; Apaéstegui et al., 2014), and the RdlP freshwater continental supply (Perez et al., 2016), related to the MCA and LIA. Modified from Perez et al. (2016).



RESEARCH PAPER

2.2 X-ray fluorescence sediment core scanning

Relative intensity distributions of major chemical elements (measured as counts per second, cps) were obtained by XRF sediment core scanning (AVAATECH core scanner at MARUM, University of Bremen) at 1 cm resolution. To remove the dilution effects, element ratios were plotted following the approach described by Weltje and Tjallingii (2008).

In the present study, we re-assessed Fe/K, Ti/Al, Fe/Ca and Ti/Ca as regional proxies for cyclicity in continental runoff and fluvial supply to the inner Uruguayan continental shelf, as previously used in Perez et al. (2016). Also, we particularly tested the use of Si/Al as a proxy for the reconstruction of marine primary productivity.

2.3 Fourier transform infrared spectroscopy (FTIR)

FTIR analysis was performed for the estimation of biogenic silica (BSi) content. Sub-samples were taken every 3 cm throughout the whole sediment core, except for the upper 100 cm where the material was sampled every 1 cm. The samples were dried at 100 °C and ground with an agate mortar and a pestle. An aliquot of 0.015 g of the sample material was mixed with 0.15 g of potassium bromide (KBr, which was dried at 105 °C for 2 h before use and then placed in a desiccator overnight). The mixed powder was pressed using a 10-ton hydraulic press to generate a transparent pellet, which was stored in a desiccator for a few hours before analysis. A pure KBr pellet was analyzed as background blank before sample analysis.



Fig. 3. (a) Study area. The red circle indicates the location of sediment Core GeoB13813-4 retrieved from the inner-shelf mud depocenter off the Uruguayan coast (map modified from Freplata, 2004). (b) PARASOUND sub-bottom echosounder profile showing the stratified mud depocenter inside the RdIP paleo-valley and its sedimentary filling succession (from Lantzsch et al., 2014). (c) GeoB13813-4 core lithology (from Krastel et al., 2012). Stars on the right of the sediment core indicate the depth of radiocarbon ages. The red arrow indicates the RdIP mud depocenter and yellow line corresponds to the cross-section sub-bottom profile. Extracted from Perez et al. (2016).



A Frontier FTIR spectrometer (Perkin Elmer, USA) was utilized using the KBr technique (Petrovskii et al., 2016). The instrument operated at a range of wavelengths between 400 and 4000 cm⁻¹ and in the absorbance mode. Twenty scans were averaged at a resolution of 2 cm⁻¹. Spectra were stored and analyzed using the Perkin Elmer spectrum software version 10.03.06.0100. Data treatment was performed using the Origin Pro software (OriginLab, Northampton, MA). FTIR raw data were smoothed using the Saviztky-Golay algorithm with a second-order polynomial and 7 points window. The baseline was carefully determined and subtracted afterward. Data were normalized to the highest peak at 1030 cm⁻¹. Considering that the band at 800 cm⁻¹ is the most appropriate for estimation of BSi content (Petrovskii et al., 2016), we selected the region of 830-660 cm⁻¹ for the partial least squares (PLS) regression (Lindberg et al., 1983; Wold et al., 2001). In order to perform PLS with full cross-validation, a calibration model was constructed using pure diatomaceous powder (Sigma Aldrich) as a BSi standard signal. A synthetic sediment mixture was prepared by adding BSi in consecutive steps to a reference sediment sample, which was previously screened to ensure a low content of BSi. Such a progressivelyincreasing mixture of a sediment sample with diatomaceous powder was thoroughly homogenized in a mortar and subsequently analyzed by FTIR (Meyer-Jacob et al., 2014). The optimum number of terms in the PLS calibration model was selected by minimizing the standard error of crossvalidation.

2.4. Statistical Analyses

The continuous time series of element ratios were analyzed with the PAST software (Hammer et al., 2001). First, we performed a spectral analysis on the element ratio series (Fe/K, Ti/Al, Fe/Ca, Ti/Ca and Si/Al) following the REDFIT procedure (Schulz and Mudelsee, 2002). REDFIT allows direct processing of unevenly spaced time series, by fitting a first-order autoregressive (AR1) process, characteristic of many climatic processes (Schulz and Mudelsee, 2002). With the REDFIT procedure, the interpolation in the time domain and its inevitable bias can be avoided (Schulz and Mudelsee, 2002). The estimated AR1 model is transformed from the time domain into the

frequency domain. The spectrum of the irregularly spaced time series is determined without the need for interpolation using the Lomb-Scargle Fourier transform (Lomb, 1976; Scargle, 1982, 1989). When performing the spectral analysis with REDFIT, we considered the Welch overlapped segment averaging estimator (Welch, 1967) by dividing the time series into five sections with 50% overlap, and by fitting within each segment, the Lomb-Scargle periodograms considering a Welch window. The resulting periodograms were averaged to obtain the final estimate. To detect significant peaks, the spectrum of a reference red noise was calculated within a 95% confidence interval assuming a χ^2 distribution, (Schulz and Mudelsee, 2002). Peaks exceeding the upper limit of the confidence interval were considered as significant (Schulz and Mudelsee, 2002). The individual periods of the significant peaks were calculated using the equation: 1/peak frequency.

This spectral analysis indicates stationary periodicities, which means they do not change dramatically in either amplitude or frequency through time. In many cases, however, non-stationary periodicities are of interest as well (Hammer, 2010). Morlet wavelet analysis is a suitable tool for analyzing these localized variations of power in a time series. By decomposing a time series into time-frequency space, it is possible to determine both dominant modes of variability as well as mode variability with time (Torrence and Compo, 1998). However, the Morelet wavelet analyses do not allow direct processing of unevenly spaced time series but constant times. Thus, based on the striking linear trend observed in the age-depth model and the homogeneous lithology throughout the whole sediment core GeoB13813-4, we assumed a mean linear constant net sediment deposition rate of 1.1 cm yr.-1 for wavelet analyses. Based on a mean historical sedimentation rate of 1.1 cm yr.-1, each 1 cm sampling interval was equivalent to a step of 0.9 yr.

3. Results

The sedimentation rates between dated intervals ranged from 0.8 to 1.3 cm yr⁻¹, from bottom to top of the sediment core, with fairly constant values of 1.0-1.1 cm yr⁻¹ from 305 to 705 cm depth (Tab. 1). The slope of the linear model generated with the weighted average age was of 1.1 cm yr.⁻¹. The base of the core was dated to 908 cal yr BP (Fig. 4).

Tab.	1.	Radioca	arbon	dates as	obtained	from	the	Bacon	modeling.
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Lab code # (Poz-)	Depth (cm)	Material	Raw ¹⁴ C AMS age (yr BP)	2 σ (95%) calibrated age (cal. yr BP)	Sedimentation rate (cm yr ⁻¹)
35198	255	bivalve shells	640± 30	62.9-211.8	1.3
47935	305	bivalve shells	775± 35	102.4-249.7	1.1
42428	447	bivalve shells	1000 ± 40	247.7-388.4	1.0
35199	560	bivalve shells	1090 ± 30	347.5-490.3	1.1
47937	705	bivalve shells	1220± 40	484.4-654	1.0
42429	964	bivalve shells	1600 ± 30	786.7-1102	0.80



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Fig. 4. The age-depth model for Core GeoB13813-4 using the Bacon program. Upper panels depict the Markov Chain Monte Carlo (MCMC) iterations (left), the prior (green curves) and posterior (grey histograms) distributions for the sedimentation rate (middle panel) and memory (right panel). The bottom panel shows the calibrated ¹⁴C dates (transparent blue), extraction year of the core (-59 yr. BP, 2009 AD, transparent blue light) and the age-depth model (grey stippled lines indicate the 95% confidence intervals; the red curve shows the 'best' fit based on the weighted mean age for each depth). Sedimentation rates are shown in green.

The elemental data showed similar vertical distributions throughout the core depth (Fig. 5). In all cases, lowest element levels were observed from 600 cm to the bottom of the sediment core, while higher and more variable values were detected from 500 cm to the top. Ti/Al, Fe/K, and Si/Al ratios displayed a succession of distinct maximum/minimum peaks from 400 cal yrs BP to present (Fig. 5). All element ratios showed evident increases during the last century, particularly over the past 50 years (Fig. 5).

The spectral analysis of the element ratios performed with the REDFIT method, revealed various significant cycles of multidecadal and centennial variability (i.e. 215, 192, 115, 100, 49, 47, 37, 35 yrs cycles) but also cycles of shorter duration (i.e. at interannual scales ranging from 2.5 to 8 yrs; Fig. 5). Fe/Ca and Ti/Ca presented a comparable pattern with significant cycles of 215 yrs (Fig. 5). Furthermore, Ti/Ca and Fe/Ca showed cycles of 2.5-5 yr but this were too close to 0 and they do not appear as significant in the wavelet analysis (Fig. 5). Spectral analysis of Ti/Al displayed significant cycles of 47, 37 and 2.7-8 yr, while Fe/K and Si/Al displayed similar cycles of 215, 115, 49 and 2.5-5.2 yrs, and 192, 100, 47, 35 and 2.7-3.7 yrs, respectively (Fig. 5). The wavelet analyses of all five element ratios showed that the 192 and 215 yrs cycles were very robust and persisted throughout the whole sediment core, i.e., over the past 1,000 yr BP (Fig. 5). The 35, 37, 47, 49, 100 and 115 yrs cycles were more significant for the Ti/Al, Fe/K and Si/Al and more pronounced during both the LIA and contemporaneous times than during the MCA (Fig. 5). The interannual 2.5-8 yrs cycle was only significant in the wavelet analysis for Ti/Al and Si/Al during some periods of time within the last 400 yr BP (Fig. 5).

The FTIR analyses revealed that the biogenic silica did not show a trend comparable to that of the Si/Al curve (Fig. 6) and the correlation between Si/Al element ratio and biogenic Si was low ($r^2 = -0.11$; p < 0.05).

4. Discussion

4.1. Late Holocene RdlP hydroclimatic history

The XRF element ratios must be interpreted as part of an environmental scenario with a high supply of terrigenous continental sediments (source) deposited in the inner Uruguayan continental shelf (Lantzsch et al., 2014; Perez et al., 2016; Mourelle, this volume). Changes in terrigenous supply are associated with the variability in freshwater discharge to the inner shelf and, therefore, to the variability in rainfall over the RdlP watershed. The geochemical and biological evidences throughout the sediment core (δ^{13} C and C/N values ranging between -20 ‰ and -18.8 ‰ and 9 and 12, respectively, and both influence of freshwater and diatom marine salinity-indicative groups and palynomorphs), suggest a mixed fluvio-marine signal within the RdlP mud depocenter (Lantzsch et al., 2014; Perez et al., 2016; 2017; Mourelle, this volume).



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Fig. 5. Time series analysis for Ti/Al, Fe/K, Si/Al, Fe/Ca and Ti/Ca. Element ratio profiles are shown in the left panel, the Lomb periodograms using the REDFIT algorithm are shown to the right. The green line in the periodogram represents the 95% confidence interval assuming a χ^2 distribution. Significant peaks are labeled with their corresponding periods in yr. The results of the wavelet analysis, are shown in the center panel. The abscissa axis corresponds to the depth of the sediment core. The ordinate axis represents the period in yr. Black lines indicate the 95% significance level and the cone of influence (COI). Inside the COI, the edge effects may become important and the results should be considered with precaution (Torrence and Compo, 1998). The color scale represents the spectral power. Major climatic changes are indicated with dotted lines, i.e., MCA and LIA.

Perez et al. (2016; 2017) identified two main stages, representing two contrasting salinity regimes over the inner continental shelf adjacent to the RdlP, during the last 1000 yr BP. The first, concordant with the MCA, characterized by a reduction in both SAMS and SACZ activities, causing a decrease in the rainfall rate, thus resulting in estuarinemarine conditions predominating over the freshwater plume signal. During the second, i.e., the LIA, a strengthening in both SAMS and SACZ activities, should have led to increased rainfall, and therefore, stronger terrigenous/freshwater supply to the inner shelf.

4.2. Element ratios and paleoclimate implication

Numerous studies have used major elements in marine sediments to reconstruct climatic history, but the choice of particular element ratios and the interpretation of such proxies, vary from site to site (Govin et al., 2012). The composition of the terrigenous material that determines the element ratios depends not only on its source (i.e., rock and soil types and their degree of chemical weathering), but are also affected by the input of biogenic opal, calcite, and mafic rock material (Govin et al., 2012). Regarding the sediments influenced by the RdlP river discharge, which are not subject to aeolian input (Mahowald et al., 2006), Govin et al. (2012) observed that Ti/Al, Si/Al and Fe/K ratios resemble the degree of chemical weathering within the drainage basin. This is related to the fact that terrigenous suspended material from the RdlP drainage basin, is mainly related to a mafic rock origin from the arid and mountainous Bermejo River sediments, which are rich in Ti, Si and Fe (Depetris and Griffin 1968; Depetris et al., 2003; Govin et al., 2012). However, biogenic Si can influence the Si/Al ratio (Govin et al., 2012). Hence, to test the use of Si/Al as a proxy for productivity in the study area, we measured the biogenic silica content and calculated the correlation between BSi and the Si/Al ratio, but we observed a weak negative correlation. Thus, we interpreted the Si/Al element ratio, as a proxy for the intensity of continental chemical weathering, also used by Chiessi et al. (2010) for the RdlP watershed. Furthermore, Ti/Ca and Fe/Ca was highly used as proxy for the RdlP continental supply versus the SWAO influence (i.e., Chiessi et al., 2009; Mahiques et al., 2009; Govin et al., 2012; Bender et al., 2013; Burone et al., 2013). Therefore, in this study, we propose that Fe/K, Si/Al, and Ti/Al truly reflect the terrestrial climatic history over the RdlP drainage basin during the Late Holocene. The variation of such ratios should be related to the changes in the precipitation patterns, which determine the drainage basin chemical weathering, and thus the continental runoff into the adjacent SWAO. In contrast, Ti/Ca, and Fe/Ca do not allow to reliably infer terrestrial climatic conditions as there is a proportion of terrigenous supply relative to the marine calcite production, and thus these element ratios are affected by dilution (Govin et al., 2012). The element ratio cyclicity was determined using time series analysis, and the observed high sedimentation rate, allowed us to infer not only centennial variability for all five ratios, but also further significant cycles of multidecadal and interannual variability (for Fe/K, Si/Al, and Ti/Al).

4.3 Centennial and multidecadal cycles

The wavelet analyses of all five element ratios indicate that the 215 and 192 yr cycles, registered for all periodograms, are very robust, and persistent over the entire record. Such cycles can be related to solar forcing (Patterson et al., 2007; Galloway et al., 2013), e.g., related to the De Vries/Suess 200 yr cycle of solar activity (Novello et al., 2012; Lüdecke et al., 2015). In this sense, such 200 yr Vries/Suess cycle, appears to be the major driver affecting SAMS intensity at centennial time scales, impacting on the northeastern Brazilian rainfalls (Novello et al., 2012, 2016). During warming in the northern hemisphere because of increased solar irradiance, there is a northward shift of the ITCZ with a concomitant weakening of the SAMS (i.e., MCA; Bird et al., 2011; Apaéstegui et al., 2014). Thus, a decreasing RdlP river discharge was inferred during the MCA, while the opposite pattern was inferred during periods of reduced solar irradiance, i.e., LIA (Perez et al., 2016).

The 115, 100, 49 and 47 yr cycles inferred from Ti/Al, Si/Al, and Fe/K element ratios are most probably related to the AMO (Fig. 5; Deser et al., 2010; Ólafsdóttir et al., 2013). Such climatic oscillation cycles were reported for SESA by Chiessi et al. (2009), Seager et al. (2010) and Apaéstegui et al. (2014). AMO is an ongoing series of long-term changes in SST in the North Atlantic Ocean and is considered to be a natural oscillation mode of the Atlantic Ocean's thermohaline circulation (Delworth and Mann, 2000; Deser et al., 2010). AMO exerts an effect on SAMS and SACZ intensity, leading to reduced intensity during positive AMO phases, thus decreasing the precipitations over SESA and vice versa (Chiessi et al., 2009; Apaéstegui et al., 2014). Consequently, a decreased RdlP river flow is expected during



RESEARCH PAPER

times of lower precipitations (Tab 2). Furthermore, it was documented that during a North Atlantic cooling (negative phase of AMO) a southward shift of the ITCZ takes place, accompanied by an associated increasing of Southern Ocean mid-latitude surface westerlies (Lee et al., 2011). Therefore, an increase in the RdlP river discharge is expected together with the dominance of SW winds, which pushes the RdlP plume northwards (Piola et al., 2008). Despite the multidecadal cycles are evident in Ti/Al, Si/Al and Fe/K periodograms, the intensity of the cycles of all five element ratios, throughout the time, is higher during the LIA and contemporaneous times comparing with those of the MCA (Fig.5). Such a pattern is particularly distinct in Ti/Al, Si/Al and Fe/K wavelet analysis plots, and can be related to the fact that these terrigenous element ratios truly reflect the continental climatic changes over the RdlP drainage basin. On the other hand, Ti/Ca and Fe/Ca are affected by dilution (Govin et al., 2012). Furthermore, this issue explains the absence of such a multidecadal cyclicity in the Fe/Ca and Ti/Ca periodograms.



Fig. 6. Variation in Si/Al and biogenic silica (%) in sediment Core GeoB13813-4 for the last 1,000 cal yr BP.

4.4 Decadal and interannual cycles

The shorter frequency cycles (2.5-8 yr cycles) registered only for Ti/Al, Fe/K and Si/Al ratios are most probably related to the ENSO mode of variability (Deser et al., 2010). Similar ENSO short oscillation cycles were determined for both Uruguay and Paraná River (Robertson and Mechoso,

1998; Krepper et al., 2003). Robertson and Mechoso (2000) and Seager et al. (2010) demonstrated that El Niño conditions have led to increased river flows of SESA. Furthermore, the significant decadal 35 and 37 yr cycles, registered for Ti/Al and Si/Al ratios, are most probably associated with PDO (Deser et al., 2010; Fig.5). PDO and ENSO are related to changes in the Pacific SST (Delworth and Mann, 2000; Deser et al., 2010). ENSO is an interannual anomaly from expected SSTs in the equatorial Pacific Ocean, while PDO is described as a decadal El Niño-like pattern of the Pacific Ocean climate variability (Garreaud et al., 2009). PDO is, at the same time, associated with ENSO as both seem to produce similar climatic effects, although their intrinsic mechanisms are not yet fully understood (Garreaud et al., 2009). Such climatic oscillations play an important role in the precipitation variability within the RdlP watershed (Tab 2; Chiessi et al., 2009; Garreaud et al., 2009). It has been suggested that during both warm phases, there is an increasing trend in precipitation over the RdlP watershed, associated to an intensification of the SAMS and higher RdlP discharge, while the opposite trend was inferred for the cold phases (Tab 2; Ciotti et al., 1995; Depetris and Pasquini, 2007 b; Garreaud et al., 2009; Barreiro, 2010; Marrero et al., 2014). Thus, we observed higher values of Ti/Al, Si/Al and Fe/K during LIA which is considered to be an El Niño-like anomaly, while we recorded the opposite pattern during MCA (Pausata et al., 2015).

4.4 The Atlantic and Pacific teleconnections

The AMO was shown to modify the characteristics of interannual variability in the Pacific (i.e., ENSO), via atmospheric teleconnection patterns, where changes originating in the Atlantic appear to exert an effect on the base-state properties of the Pacific, by altering the trade wind characteristics (Delworth et al., 2007). Therefore, a warm phase of the AMO leads to a weakening in ENSO variability. In contrast, when the northern hemisphere cooling shifts the ITCZ southwards, it triggers a weakening in trade winds over the western and central equatorial Pacific, which propitiates the development of an El Niño-like anomaly (e.g., during LIA; Pausata et al., 2015). This is the most plausible explanation for our observation of an intense AMO followed by more intense PDO and ENSO during the last 500 yr BP. Thus, we infer an increase in the fluvial discharge and terrigenous material supply during the LIA and contemporaneous time due to an intensification of AMO, PDO and ENSO signal. The reason why we only registered AMO, ENSO and PDO climatic modes of oscillations within the Ti/Al, Fe/K and Si/Al element ratio, is because they truly reflect the terrestrial climatic history over the RdlP drainage, while Ti/Ca and Fe/Ca are related to marine productivity also modulated by oceanographic processes. This study introduces further evidence that both tropical Pacific and Atlantic SST and solar forcing are the



main drivers of SESA changes in rainfall at interannualdecadal, multidecadal and centennial scales. Also, the RdlP watershed precipitation strongly depends on such Pacific-Atlantic interactions (Seager et al., 2010).

5. Conclusions

Changes in the XRF element ratios studied from a sediment core retrieved from the RdlP mud depocenter mainly follows the paleoclimatic history of the RdlP drainage basin, but also the paleoceanographic history of the SWAO (related to the marine productivity). In this sense, Ti/Al, Fe/K and Si/Al, follows the degree of chemical weathering and thus, record the terrestrial climatic history over the RdlP drainage basin during the Late Holocene, mostly related to changes of the precipitation patterns. Furthermore, Ti/Ca, and Fe/Ca reflect the continental supply versus the marine biogenic calcite production in the RdlP estuary and the coastal SWAO region and thus is affected by dilution. Thus, Ti/Al, Fe/K and Si/Al were more informative in terms of decipher the climatic oscillation modes affecting the continental input into the Uruguayan inner continental shelf.

This study shows that the terrigenous sediment supply from the RdlP to the continental shelf was strongly controlled by long-term regional climatic cyclicities: i.e., 215 and 192 yrs cycles, most probably related to solar forcing (i.e., the 200 yrs Vries/Suess cycle), as well as by shorter climatic cyclicities (<100 yrs). The most significant shorter cyclicities occur at 115, 100, 49, 47, 37, 35 and 2.5-8 recurrence times, only for Ti/Al, Fe/K and Si/Al. Such shorter climate-related cyclicities are most probably related to the AMO, the PDO and the ENSO climatic modes of oscillations.

Furthermore, the variability of these competing climatic modes modulates the latitudinal migration pattern and the regional establishment of ITCZ, SAMS, and SACZ. As a consequence, these climatic modes and their variability determine the regional precipitation patterns and weathering over SESA and the RdlP drainage basin, and thus determine the input of terrigenous material, as inferred from the element ratios of the inner Uruguayan continental shelf sediments. Furthermore, we observed an intensification of the multidecadal, decadal and interannual cycles during the last 500 yr BP.

Tab. 2. Climatic oscillation modes and their relationship to rainfall in the RdlP watershed.

	Climatic oscillation modes							
	AMO		PDO		ENSO			
Phases	+	-	+	-	+	-		
RdlP Watershed Rainfall	Decrease	Increase	Increase	Decrease	Increase	Decrease		
Reference	Chiessi et al. (2009); Seager et al. (2010)		Garreaud et al. (2009)		Depetris and Pasquini (2007b); Garreaud et			
Reference					al. (2009); Seager et al. (2010)			

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