ALKALINE COMPLEXES AS POTENTIAL AREAS FOR INSTALLATION OF REPOSITORIES OF NUCLEAR WASTE: CRETACEOUS INTRUSIONS IN BRAZIL

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Abstract

Brazil initiated its nuclear program in the decade of 70 with the installation of the Thermonuclear Plant Angra I. The implantation of this type of enterprise involves activities which generate environmental risks mainly due to disposal of nuclear waste. The storage of nuclear waste must follow strict international standards and faces the resistance of the civil society. Thus, complex technical and politically procedures are required to find a suitable place for this end. The use of geologic bodies that fit the required features for the installation of nuclear repositories is described in literature. This kind of geological units must be deep, structuralized and with low infiltration. These aspects should to be identified through geologic, hydrologic and geophysical studies and structural models as well.

This study aims to report the results of the regional geophysical survey performed to delimit the gneissic host rocks of Ribeira Belt and the characterization of their magnetic characteristics main textural variations. This work also includes the analysis of the structural features of rocks (faults and fractures), in addition to how these structures change within the massif coupled with mineralogy and chemistry signatures.

The magnetic anomaly map displays a series of semi-circular magnetic highs, following the general configuration of the alkaline intrusions. In Brazil, the intrusive bodies with potential to function as repository are the alkaline complexes, once they are isotropic, massive, have well defined geometry, compatible depth and with low or no human occupation. However special attention to the possible occurrence of water percolations in fractures and faults must be given. In this way, detailed studies of alkaline bodies that fit the expected features comprise the following intrusions: Mendanha, Tinguá, Tanguá, Soarinho, Rio Bonito and Morro de São João. This study intends to be a contribution for the knowledge of the properties of alkaline bodies with potential to be used for storage of nuclear waste, in the near future, in Brazil.

Keywords: Nuclear waste. Repository. Radioactive waste management. Alkaline complexes.

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1. Introduction

The disposal of radioactive waste is an issue relevant for the nuclear industry in Brazil. At present, there is no knowledge about its final destination, especially the waste from the spent fuel from nuclear plants of Angra dos Reis, classified as having high radioactivity. In terms of nuclear waste of high radioactivity, the literature emphasizes that geological repositories are ideal places to house these types of materials. Documents divulged by the International Atomic Energy Agency (IAEA), are listed below among which:

- Site Selection Factors for Repositories of Solid High-Level and Alpha Bearing Wastes in Geological Formations (IAEA, 1977);
- Underground Disposal of Radioactive Wastes (IAEA, 1981);
- Criteria for Underground Disposal of Solid Radioactive Wastes (IAEA, 1983a);
- Concepts and Examples of Safety Analyses for Radioactive Waste Repositories in Continental Geological Formations (IAEA, 1983b; IAEA, 1994a);
- Regulatory Objectives, Requirements and Guidelines for the Disposal of Radioactive Wastes, Long-Term Aspects (IAEA, 2009);
- Guidance for Regulation of Underground Repositories for Disposal of Radioactive Wastes (IAEA, 1989);
- Siting of Geological Disposal Facilities (IAEA, 1994b);
- Geological disposal of radioactive waste (IAEA, 2011);

Based on such perspectives, Enokihara (1983) evaluated that geological formations more favorable for the deposition of nuclear waste are classified in four basic groups: evaporites, other sedimentary rocks, stones and igneous rocks.

As regards the lithologies, especially in relation to the long-term safety associated with nuclear waste, Martins (2009) underlined that the insertion crystalline areas, mainly composed by metamorphic rocks and magmatic rocks (among which alkaline, granite and basic rocks), are the most propitious repositories of nuclear waste. However, these rocks vary enormously in strength, stability, homogeneity, chemical composition, among other factors. This is why not all rocks classified in this way are adequate for the construction of a repository. The greatest problems regarding its use for such end is the presence of failures and/or fractures that weaken some rocky areas and may constitute paths of migration of radionuclides.

The site description is basically needed aiming to: (1) develop and present the geoscientific and ecological understanding of the site; (2) underpin a safety assessment describing different scenarios and consequences if radionuclides are released from the repository; (3) establish a baseline for detecting short- and long-term effects of the repository; (4) provide the basis for the environmental impact assessment describing and evaluating the effects of the construction and operation of a repository; (5) provide input to the development of preliminary designs of the surface and subsurface constructions.

1.1. Objectives

In this context, the present work aims to identify and characterize the alkaline complexes as potential areas for future radioactive nuclear waste repositories, which include alkaline complexes located in the southeastern region of Brazil. In the Rio de Janeiro State and its adjacent areas, there are a dozen of felsic alkaline intrusive rock bodies forming WNW-ESE alignments (Amaral et al., 1967; Almeida, 1976), called Poços de Caldas - Cabo Frio alkaline rock alignment (Figure 1), such as Itatiaia, Morro Redondo, Tinguá, Mendanha, Itaúna, Tanguá, Rio Bonito, Soarinho, Morro dos Gatos, Morro de São João, Cabo Frio Island (Motoki, 1979; Brotzu et al., 1989, 2007; Mota et al., 2009; Motoki et al., 2012; Geraldes et al., 2013).

Initially, we discuss about the issue of radioactive waste, presenting their main characteristics. Subsequently, we emphasize the importance of alkaline rocks as repository, especially those with high radioactivity. This contribution has as objective to subsidize the final decision on the location and construction of repositories for the final deposition of radioactive waste in the aspects of site selection, characterization of local, development of technologies for optimizing the performance of the host rocks deposition in the system, the security analysis in function of the underground aquifer, and support to nuclear and environmental licensing.

The operation of the nuclear power plants in Brazil and the use of radioisotopes generate considerable quantities that require the deployment of plants licensed for the deposition, making urgent the definition of a concept of repository for the country, for their subsequent construction and operation. In this sense, the urgent requirement is the beginning of studies for the site selection, design, construction, and licensing for the entry into operation in the shortest time.
1.2. Radioactive Waste issues

In agreement with the IAEA (2009), the radioactive wastes are classified into six categories as a function of its activity, as shown in Table 1. The radioactive waste produced in the Angra dos Reis Nuclear Power Plant may be classified in three levels of radioactivity: high, medium and low (IAEA, 2009). For the group of low and average radioactivity (which should stay in the middle tank by a maximum of three years) the destination are two concrete sheds constructed within rocks, beside the plant, where it is stored in drums containing boots, overalls, contaminated clothing used by workers or the metal parts of the reactor and chemical waste.

The group of high radioactivity are formed by already irradiated fuel element inside the reactor, whose half-life is quite long, and may reach tens of thousands of years, making the question on the fate of these wastes of high relevance. Currently, the place of storage of these wastes are done in pools inside the reactor area that keeps the waste submerged at more than 10 meters deep (Freire and Tello, 2007; 2008). However, this type of reject, with the decommissioning of the nuclear power plants, must have as final destination a geological repository.

1.3. Geological Repositories of Radioactive Rejects

As underlined by Silva et al. (2015), according to the IAEA the repositories for radioactive waste can be classified in two categories: (A) next to the surface repositories and; (b) geological repositories. The repositories close to the surface are recommended for the radioactive waste of low and middle risk, being then the establishment of an institutional control period is quite important, in view of the risk of intrusion in the long term. The geological repositories (Neves et al., 2007; Silva and Tello, 2007), applicable to high-level radioactive waste, unlike deposits coming to the surface, are located at great depths, aiming to eliminate the risk of intrusion of humans and animals.

Raduan (1994) evaluated the geological formations as the most adequate to the insulation of dam high level or tailings containing radionuclides with long half-life, radionuclides Transuranium Elements, among others. According to the IAEA (2009), some tailings need a deep geological disposal to avoid problems for the environment and for society. Among such waste, we can highlight those that participate of the nuclear fuel cycle, including own fuel used in the process of generating energy. In addition, the security of a geological deposit is based on the concepts of defense in depth and in isolation, which is supplied by the combined effects of multiple artificial and natural barriers.

In addition to the properties of the host rock, the specific aspects of the project and the degree of filling and sealing of the openings and connections with the surface are highly relevant. Furthermore, certain characteristics of the project, such as the containment of long-lived waste, are clearly...
beneficial. Whereas the geological aspects and its importance as a deposit of radioactive waste, the IAEA observed that hard rocks (granites, for example) are effectively self-sufficient, being necessary only small engineering and maintenance measures to prevent the bank rupture of the walls of the rock in cells of deposition and places of access.

Although there are several concepts on the isolation of radioactive waste in the biosphere, during the time needed for its decay, Enokihara (1983) observed that the only alternative that appear viable, from various points of view, is your storage in deep geological repositories, whose main attraction is the fact that many geological formations remained stable for millions of years. In addition, these formations present the three basic requirements for the safe disposal of radioactive materials: cooling, protection and insulation. So, the surrounding rock acts as a heat sink, providing the cooling. The depth of deposition provides adequate protection against radiation, and the size and integrity of rocky formations enables Isolation, preventing the escape of material to the biosphere.

2. Study area

2.1 Geomorphology of Geological Sites Repositories

Alkaline intrusions are constituted by stocks of the size varying from less than 1 km up to 15 km composed of nepheline syenite and alkaline syenite. Small bodies of phonolite and trachyte also occur. In Morro de São João intrusive complex (Brotzu et al., 1989; Mota et al., 2009) and Cabo Frio Island gross-grained mafic alkaline rocks are found. Lamprophyre dykes show fine granulation. These rocks are relatively common in the Rio de Janeiro State and extracted as highly qualified as ornamental construction material (Petrakis et al., 2009). Some of the intrusive complexes such as Itatiaia, Mendanha (Motoki et al., 2012), Itaúna (Motoki et al., 2013a), Morro dos Gatos and Cabo Frio Island (Geraldes et al., 2013) have subvolcanic pyroclastic conduit and fissures made up of pyroclastic rocks (Mota and Geraldes, 2006; Motoki et al., 2013b). Along the alignment, there is Cambrian nepheline syenite gneiss of different genesis called Canaí complex. The intrusive bodies have circular (e.g. Tanguá, Morro de São João) or elliptic (Mendanha, Rio Bonito, and Cabo Frio Island) distribution. The long axis of the elliptic bodies ranges from ENE-SWS to NE-SW.

The main intrusive bodies have rare homogeneous form (Geraldes et al., 2013). That is, the horizontal section of upper level of the pluton is circular and that of lower level is elliptic. Their tilling is supposed to be sub-horizontal and the bottom is considered to be fissure-like feeder of the felsic alkaline magma.

The nepheline syenite and alkaline syenite is made up mainly of alkaline feldspar, nepheline, alkaline clinopyroxene, amphibole, and biotite with accessory minerals of ilmenite, magnetite, apatite, and titanite. The alkaline feldspar can show interlocking perthite but no flame perthite and string perthite are observed. The perthite type indicates that this feldspar it is not so strongly potassic but intermediate one. Several rock bodies, such as Tanguá has amphibole but no clinopyroxene (Motoki et al., 2012). The Morro dos Gatos syenite has exceptional composition on the border line between syenite and monzonite. This rock contains clinopyroxene and biotite but no amphibole is observed. Zircon is not found in nepheline rich rocks. The amphibole occurs frequently along the border of the clinopyroxene and some amphibole crystals are pseudomorph of clinopyroxene. The rocks without nepheline can have little amount of zircon and quartz. The nepheline syenite of Tanguá and Morro de São Joao complexes present local occurrence of pseudoleucite.

The alkaline intrusive complexes (nepheline syenite, alkaline syenite, phonolite and trachyte) form morphologic elevation with relative height of 300 to 900 m, called locally alkaline massifs. Their minerals (especially nepheline and alkaline feldspar) are vulnerable to chemical weathering under humid tropical climate. This controversy is explained by the phenomenon called weathering passivity (Motoki, 1979), that is, the highly impermeable regolith composed of clay minerals originated from the chemical weathering protects the massif preventing the surface water percolation into the rock body. Therefore, the alkaline rocks present just below the regolith is fresh and mechanically very firm.

According to the fission track dating for apatite, the present exposure is considered to be subvolcanic structure of 3 km below the surface of the eruption time (Figure 2). The volcanic edifices and surface eruptive materials of that time have been completely eliminated by the regional uplift of 3 km that occurred up to 40 Ma (Geraldes et al., 2013).

2.2. Origin of Cretaceous Alkaline Rocks

The origin of the felsic alkaline magmas is not clear. In outcrops, highly felsic nepheline syenite and alkaline syenite are found with rare occurrences of mafic alkaline rocks. And intermediate alkaline rocks aren't common. It is difficult to accept the fractionation crystallization model of mafic alkaline magma. In addition, there are small number of rocks at the contact zone of the Mendanha (Motoki et al., 2012) and Cabo Frio Island intrusions, silica saturated. And the existence of pseudoleucite in the nepheline syenite with low
silica indicates that relatively less-differentiated felsic alkaline magma is K-rich and it becomes Na-Rich by means of fractionation crystallization.

Motoki et al. (2013a), based on the chemical analyses and their petrological interpretation, proposed a magma evolution model in three-stages (Figure 3). The relatively less differentiated nepheline syenite magma has high $K_2O/(Na_2O+K_2O)$ ratio. This magma crystallized leucite and the residual magma has lower $K_2O/(Na_2O+K_2O)$ ratio (stage 1) and continues up to the arrival of residual magma composition at the cotectic curve. Leucite crystals transform into pseudoleucite during slow cooling of the magma. Examples of this stage are found in Tangauá and Morro de São João complexes. The next process (stage 2) is alkaline feldspar crystallization, where this feldspar is expected to be moderately potassic one with $KAlSi_3O_8:NaAlSi_3O_8$ proportion ranging from 0.6:0.4 to 0.7:0.3. The residual magma becomes from K-rich to Na-rich sliding on the cotectic curve (by means of the fractionation of the alkaline feldspar). These examples are commonly found and continues up to the lowest temperature point on the cotectic curve. Some samples (Cabo Frio Island, Itaúna, Itatiaia and Rio Bonito) indicate that the magma composition arrived at the point close to the lowest temperature. During this stage, the assimilation of continental crust rocks can take place, with the magma composition was pulled up from the cotectic curve to the granite composition, crossing over the thermal divide. This process is called stage 3 and form felsic alkaline rocks without nepheline and are found at Mendanha, Soarinho, Morro dos Gatos, and Cabo Frio Island.

The frequency of radiometric ages from post-breakup rock samples (described in detail below) from the South Atlantic (Figure 3) can be divided into four main classes: Upper Cretaceous (60-80 Ma), with more than 40 radiometric ages for the intervals; Miocene to Recent, with 10 to 20 ages; and the less frequent radiometric ages for the mid to late Tertiary (20-50 Ma) and mid Cretaceous (90-105 Ma). Cabo de Santo Agostinho magmatic province, onshore the Pernambuco-Paraíba coastal basins (Figure 1), registered the oldest magmatic event that post-dates the inception of oceanic crust in NE Brazil.

The magma evolution crossing over the thermal divide does not occur in normal crystallization fractionation because of the thermodynamic instability. However, the
thermal divide in this composition is not so high, about 50°C, and therefore, small magma super-reheating and/or low-grade fluid enrichment can solve it (Motoki et al., 2012). These phenomena can sap take place by new magma injection into the magma chamber, helping country rock melting. The magma super-reheating can form mafic and opaque minerals of desorption shape. The resorted clinopyroxene and magnetite are commonly found in the nepheline syenite and alkaline syenite of this area. The fluid enrichment could transform the clinopyroxene into amphibole. Such examples are observed in the nepheline syenite of Mendanha and Tanguá.

![Radiometric ages](image)

**Fig. 3.** Histogram of radiometric ages (from Motoki et al., 2013a).

The pyroclastic rocks present at Mendanha and Itaúna were considered as constituent rocks of surface volcanic deposits. Based on this idea, they were interpreted to be extinct volcanic edifices. However, recent researches (Motoki et al., 2013 a; b) revealed that these rocks are not components of pyroclastic flow or pyroclastic fall deposits, but of vent-filling welded tuff breccias, that is, subvolcanic conduits and fissures.

The contact planes between the pyroclastic bodies and their host rock are subvertical without intercalation of palaeosoil and organic materials. Therefore, the outcrops are considered to be intrusive contacts. Intrusive pyroclastic rocks are little known in the world (e.g. Ekren and Byer, 1976; Motoki, 1979; Reedman et al., 1987; Kano et al., 1996; Wolff et al., 1999). Some of the pyroclastic bodies has been described as pyroclastic dyke and are very small, less than 4 m wide (Motoki et al., 2012). A pyroclastic dyke of the Morro dos Gatos is made up of densely welded intrusive tuff.

### 2.3. Petrography and Magnetometry of Alkaline Intrusions

Geologic studies conducted by university and government agencies showed that the continental alkaline magmatism in SE Brazil (Figure 4) occurs as plugs, stocks, dykes and sills that present several trends, evidenced by geological and geophysical datasets: a general ENE direction for the dykes, and to NW/SE and E-W lineaments. These alkaline complexes most probably represent the roots of extinct Cretaceous volcanoes (Figure 2). Over time the volcanic superstructures have eroded and just the plutons remained, as is also observed in present day in Cameroon Line volcanoes (Gulf of Guinea) that have been deeply eroded.

In Southeast Brazil, between longitudes 42° and 47° W, a linear alignment of alkaline igneous intrusive complexes occurs with ages ranging from 85 to 55 Ma (Figure 3). Over a length of 500 km, they were emplaced into the northern segment of the Neoproterozoic Ribeira belt. With 800 km², Poços de Caldas is the largest alkaline complex in Brazil and one of the largest in the world. The alkaline intrusions present some residual relief bodies, in part formed by plutons of alkaline nature, lined up according to WSW–ENE trend (Figure 1). As an example, the smaller Itatiaia/Passa Quatro complexes, with peaks up to 2.800 m above sea level, represent the highest elevations on the Atlantic side of South America (Ulbrich and Gomes, 1981).

In addition, the Cananeia and Jacupiranga intrusions, located between Paraná and São Paulo states, are well known for their complex lithologies and some economic importance, as P-rich minerals are explored in carbonatites from Jacupiranga. In São Paulo, there are several alkaline complexes, the most important are the Ponte Nova, Passa Quatro intrusives and the São Sebastião Island, which is associated with three alkaline intrusions.

In Rio de Janeiro State the most important intrusions (Figure 4) are located between Serra do Mar (Ocean Range) and the coastline adjacent to the Santos and Campos basins (Almeida et al., 1976). Such alignment includes Mendanha Massif, Itatiaia, Morro Redondo, Pirai (or Tomazes ridge), Tinguá, Itaúna, Monjolos, Porto das Caixas, Tanguá, Soarinho, Rio Bonito, Morro dos Gatos, Guia de Pacobãba, Cabo Frio Island and São João plutons.
Fig. 4. Magnetics survey by São Paulo-Rio de Janeiro project (data available from CPRM, Brazilian Geological Survey).
3. Material and methods

This work presents geologic details of the intrusions and aeromagnetic regional features of some of the most important intrusions of Rio de Janeiro State, such as Mendanha Massif, Tinguá, Tanguá, Soarinho, Rio Bonito, Morro de São João (Figure 4) and Itatiaia.

The mapping allowed to delineate the contact between the alkaline intrusions and the basement rocks. The local basement is composed of Paleoproterozoic orthogneisses where biotite-gneiss and biotite-hornblende-gneiss appear. In addition, structure features were delineated from satellite images.

The mapping on the scale of detail (1: 10,000) within the alkaline bodies allowed the identification of felsic and mafic lithotypes, the complexity of the contacts as well as the distribution of the lithotypes. In this sense, the fieldwork followed by the petrographic studies allowed characterizing the alkaline rocks in the studied intrusions.

The aeromagnetic survey data used in this study covers the area of Rio de Janeiro State and was acquired by Companhia de Pesquisa de Recursos Minerais (CPRM; Geological Survey of Brazil) in 1978 and corresponds to a 150 m height aero survey, with 1 km of line spacing and 10 km control line spacing.

In order to study the crustal structures in detail, an aeromagnetic anomaly map with high resolution was produced. In this way, the magnetic anomalies map allowed a better comparison with the local tectonics and major geological features. The magnetic data was surveyed with a Geometrics G-803 proton precession magnetometer with the sensor placed on the tail of the aircraft. The airborne dataset was initially processed by CPRM (2000) and its results are leveled grids (250 m cell size) of several geophysical themes, such as Total Magnetic Intensity Anomaly, Magnetic Field Diurnal Correction and Analytic Signal.

4. Results

4.1 Mendanha

The Mendanha Massifs a young alkaline complex (80-53 Ma, K/Ar; 64 Ma Ar/Ar, according to Geraldes et al., 2013; Mota et al., 2014) inserted in a NW-SE magmatic alignment in Rio de Janeiro State (Figure 5). It is distinguished by including sub-volcanic and volcanic types, identified in other plutonic provinces in this alignment.

The Mendanha Massif comprises a non-saturated suite predominantly composed by foyaites, tinguaytes and nepheline syenites, with mineralogical and textural variation showing syenitic, trachitic lavas and breccias intercalation (Figure 6). The plutonic suite and trachitic lavas are associated with two main volcanic structures. The first one is a 2 km large conduct named Nova Iguacu volcano structure which consists of central agglomerate with bombs and abundant tuff and lapilli. The second one, the Lamego Chamber, is a smaller conduct of trachitic welded tuffs and volcanic breccias. Minor occurrences of barite and bauxite had been described in the Mendanha Massif (Mota and Geraldes, 2006; Mota et al., 2012). The record of the volcanic structures can suggest fast explosive magmatic processes possibly controlled by ascendant hot gases, with suspended fragments (fluidization) accumulated in flat magmatic chambers.

4.2 Tinguá

The Tinguá alkaline complex (120 km²) is located at north of Rio de Janeiro city. These rocks intrude into the Precambrian basement (Figure 7), which is represented by the Juiz de Fora metamorphic complex (no available ages for this intrusion).

4.3 Itatiaia

The Itatiaia massif is structurally linked to the SE Brazilian continental rift, which developed following an ENE to EW trending transcurrent system inherited from the Brasiliano (Pan African) Cycle and which was also reactivated during the breakup of Gondwana. This alkaline complex (Figure 8) is composed of syenites, pulaskites, foyaites, quartz syenites and magmatic breccias (Figure 9). The mean K-Ar age obtained on ten biotite samples from the Itatiaia rocks is 70.5 ± 3.3 Ma (Geraldes et al., 2013).

4.4 Tanguá

The Tanguá Hill is the western most intrusion associated with the Rio Bonito Massif, which is located between Rio de Janeiro and Cabo Frio (Figure 10). The alkaline massif yields radiometric ages of about 72 Ma (Geraldes et al., 2013) and includes several rock types, such as nepheline syenite, alkaline syenite, breccias and phonolites (Figure 11). The Tanguá alkaline intrusive body is composed predominantly of coarse-grained nepheline syenite of light grey macroscopic color, with alkaline feldspar, clinopyroxene, nepheline, amphibole and biotite. Magnetite, titanite, and apatite are accessory minerals. Zircon is scarce. The clinopyroxene (light green in parallel nicol) indicate aegirine-augite composition. They are frequently surrounded by amphibole, suggesting the reaction between the clinopyroxene and hydrated magma, commonly observed in the alkaline syenite and nepheline syenite of other intrusive complexes.
4.5 Soarinho

The Soarinho alkaline complex (148 km²) is close to Guanabara Bay and was emplaced into the Precambrian metamorphic complexes of Juiz de Fora, Paraiba do Sul and the Açungui Group (Figure 12). Alkali plutonic rocks (nepheline syenites) with minor phonolitic dykes are the main lithological types. The K-Ar age is 67±3 Ma (Geraldes et al., 2013), obtained on only one sample of amphibole separated from a nepheline syenite.

Alkaline syenite is similar to the nepheline syenite but without modal nepheline. Outcrops of this rock show light grey color in macroscopic and difficult to be distinguished from the nepheline syenite by naked eye observations (Figure 13). Its distribution is limited only along the contact zone. Alkaline feldspar, clinopyroxene and amphibole, with magnetite and apatite as accessory minerals composed the rock. The abundance of host rock xenoliths is remarkable and they show notable plastic deformation by the host magma heat.

4.6 Rio Bonito

The Rio Bonito Massif, with approximate direction WNW, is part of the alignment of alkaline rocks of the macro region. It is seated on gneisses pre-Precambrian shields of the Ribeira Belt (Figure 14) and is composed of light gray rocks with dark coarse to medium granular texture. The mineralogical composition of these rocks is composed of K-feldspar, orthoclase feldspatoids, amphiboles micropertitic, highlighting the nepheline, plagioclase, titanite, biotite, apatite and opaque minerals.

The studied rocks were classified as nepheline, foiaite and micro-foiaite, in their majority, besides normative pulaskite, and volcanic breach. The Massif is cut by phonolite dikes, trachite, phonolite porphyry and lamprophyre (Figure 15). Regarding the silica, they are intermediate. The classification by diagrams confirms that the nepheline cyanide field is most representative.
Fig. 6. The Mendanha rocks (A) are grained felsic syenite (B, C and D) and trachyte.

Geochemical investigations of Motoki et al. (2013b) have determined the generation of rocks in environment intra-plate. Diagrams Harker of studied rocks indicate alkaline magmatic processes of differentiation. The patterns of REE (Rare Earth Elements) corroborate with this hypothesis which indicate positive anomalies of Eu (Europium) for more differentiated rocks and negative ones for more primitive samples.

The studied rocks are of an orogenic cratonic environment. Have origin in the mantle and the processes of formation can be correlated to the action of a mantle plume or through crustal flexure, where sediment deposited on the nearby continental shelf to the west have led to the opening of faults that have reached the lithosphere. This magmatic event presents the Ar cooling age of 65.03 ± 0.70 Ma (Geraldes et al., 2013).

4.7 Morro de São João

The Morro de São João Alkaline Complex is located in the NE part of Rio de Janeiro State, near the border fault of the Campos Basin, and shows a pronounced topography near the coastline (Figure 16).

Magnetometric survey of these rocks indicates a magnetic pole inversion recorded during the crystallization of the alkaline magma.
Fig. 7. Regional aeromagnetic features of Tinguá alkaline intrusion.

The intrusive Morro de São João is characterized by the following lithologies: coarse-grained felsic syenite (K-feldspar, nepheline, hornblende, titanite and pseudoleucite), and mafic syenite (with K-feldspar, hornblende and pyroxene). Petrography (Figure 17) and litho-geochemistry data (Mota et al., 2014) suggest a bimodal suite. Felsic and mafic magma mixing and mingling textures are locally observed indicating the presence of coeval magmas, as indicated by Ar cooling age of 84 Ma (Geraldes et al., 2013).

4. Discussion and Conclusion

The Brazil initiated its nuclear program in the decade of 70, through the installation of the Mine Osamu Utsumi in Alkaline Complex of Poços de Caldas (Minas Gerais, Brazil) and installation of the Thermonuclear Plant Angra 1, in the city of Angra dos Reis (Rio de Janeiro, Brazil).

The deployment of this type of project involves some activities and potential hazards that can be summarized into: mining with production of acid drainage and particles of uranium; production of yellow cake and eventual contamination resulting from the process; activity of the mill itself, with generation of radionuclides radioactive, possible leaks, the tailings profile produced and materials used in the maintenance of the Thermonuclear Plant. In this sense, along with energy produced by this type of plant, radon and xenon gases (elements of difficult containment) and isotopes are produced exclusively by reaction within the reactors and such products as that do not occur naturally in the environment (Efremenkov, 1989).

These isotopes present high radioactivity with emission of alpha (α) particles, beta (β) particles and gamma (γ) ray energy and have a longer half-life ranging between days and millions of years, and so the disposal of this material must meet strict international standards (IAEA, 1995, 2000, 2002, 2003, 2011). The production of that high level radioactive reject increased with the installation of the Thermonuclear Plant Angra 2, located within the same complex of Angra 1.
Fig. 8. Geologic map of Itatiaia region.
Fig. 9. The Itatiaia alkaline complex is composed of syenite (A), quartz syenite (C) and magmatic breccias (B and D).
Fig. 10. Regional aeromagnetic features of Tanguá alkaline intrusion.
Fig. 11. Plastically deformed (A) and digested (B) host gneiss xenoliths in the alkaline syenite at the south-western contact zone of the Tanguá intrusive complex.

Fig. 12. Regional aeromagnetic features of Soarinho alkaline intrusion.
Fig. 13. The Soarinho rocks are classified as nepheline syenite (A and B), micro-foidait (C), foidait (D) and volcanic breccias (E). The massif is cut by phonolite dikes, trachite, phonolite porphyry (F) and lamprophyre.
Fig. 14. Regional aeromagnetic features of Rio Bonito alkaline intrusion.
Currently the waste produced has been stored inside of own reactors of these two plants, but the expansion of the Brazilian nuclear program, requires a future storage area for this type of reject. The storage of nuclear rejects also faces the resistance of civil societies, so that to define a suitable location for this purpose is a problematic issue.

Some countries, such as Japan, Sweden and France, whose energy matrix is preferably produced in thermonuclear plants, have researched and found some geological bodies that meet the metric required the installation of nuclear repositories (Neves et al., 2006, 2007; Silva and Tello, 2007). The geological survey, facing the correlation with geophysical information aimed to delimit the gneissic host rocks of Ribeira Belt and to characterize their main compositional and textural variations, mainly mineralogical components, as well the structural features of rocks, in addition to how these features are distributed within the massif. These massifs have potential to be used as nuclear waste repositories. However, further detailed studies of the alkaline bodies, such as alkaline complexes here reported, may not only provide a scientific knowledge about the geological evolution and origin, but also to contribute with knowledge base enabling the future choice of geologic repository of radioactive waste.

Fig. 15. The Rio Bonito alkaline rocks are composed of alkaline feldspar, clinopyroxene and amphibole, with magnetite and apatite as accessory minerals. (A) Alkaline gabbroic rocks (Gr) cut by apliteic dykes (Apl) are observed. (B) with syenitic aplite (Sap). (C - D): details of Alkaline gabbroic rocks and apliteic dykes, respectively.
Fig. 16. Regional aeromagnetic features of Morro de São João alkaline intrusion.
Fig. 17. The Morro de São João alkaline rocks comprises a non-saturated suite predominantly composed by tinguayte (A), trachyte (B, C and D) and nepheline syenites (E and F). It presents mineralogical and textural variation showing syenitic, trachytic lavas and breccias intercalation.
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